

MINERAL DUST EMISSION PROCESSES AND THEIR MODELING: RECENT PROGRESSES AND REMAINING CHALLENGES

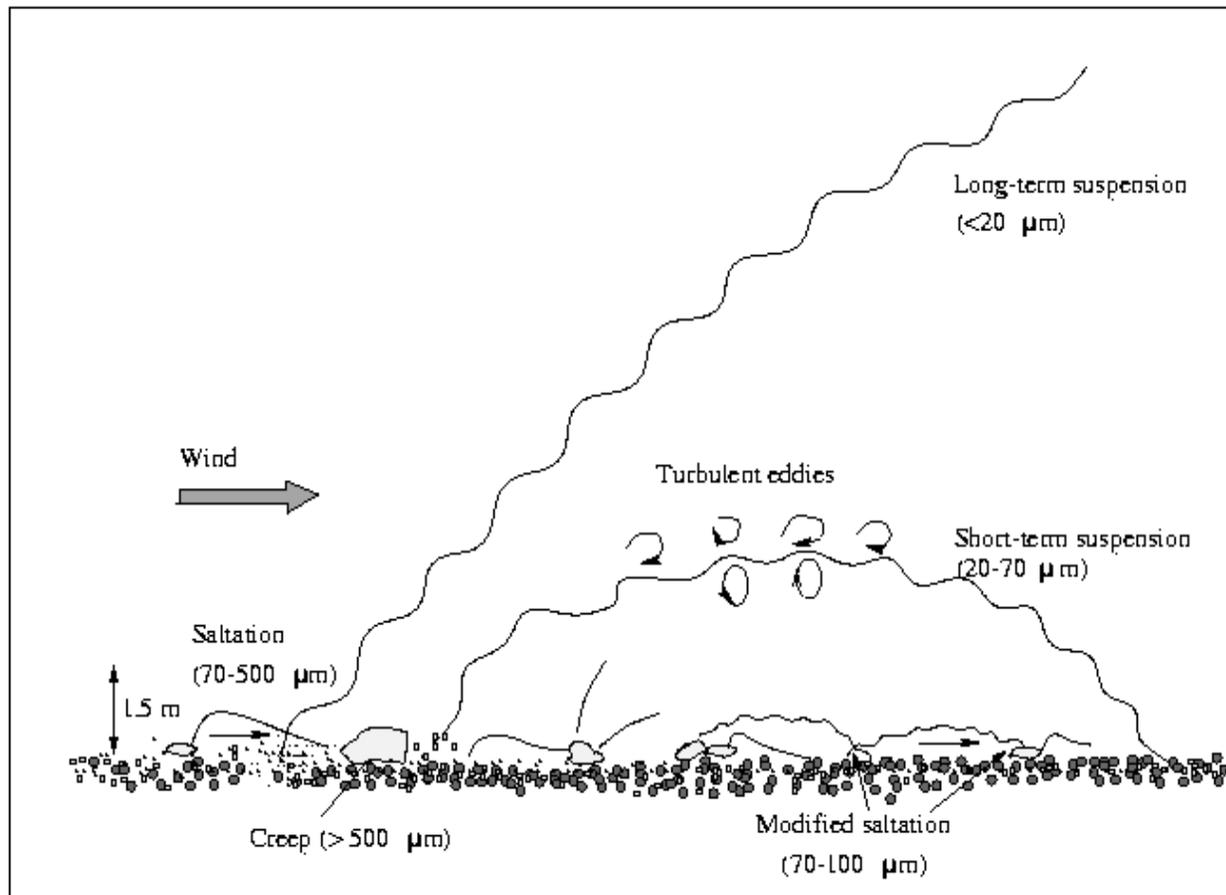
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DUST EMISSIONS PROCESSES

General picture



(From Y. Shao)

DUST EMISSIONS PROCESSES

- ✘ Dust emission results from mechanical processes involved in wind erosion
 - + **Wind is the main driver**
 - + There is a wind velocity **threshold**
 - + **Saltation** is generally a pre-requisite
 - + Dust is released by « **sand-blasting** »

$$F = \alpha \cdot C \cdot U_*^3 (U_* - U_{*t})$$

Sand-blasting efficiency \nearrow α

$\underbrace{C \cdot U_*^3 (U_* - U_{*t})}_{\text{Saltation flux}}$

\leftarrow **Threshold** U_{*t}

DUST EMISSIONS PROCESSES

Emission processes

Erosion threshold



Saltation



Sand-blasting

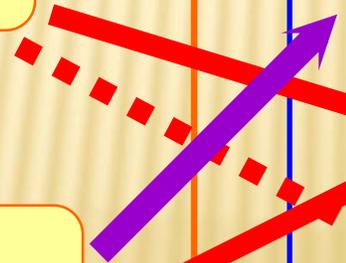
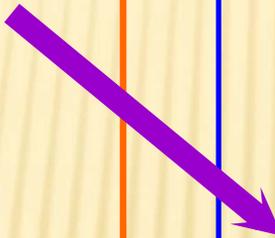
Model outputs

Location and periods

Emission flux intensity

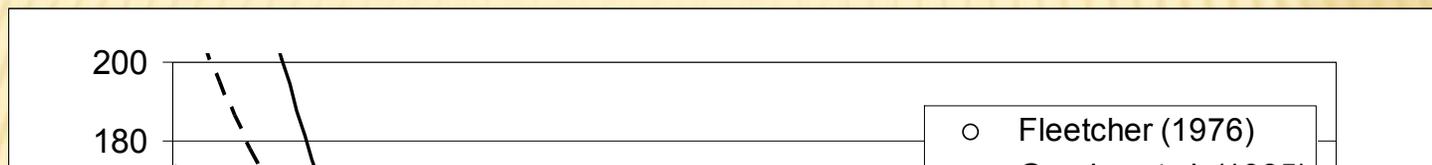
Size-distribution

Composition

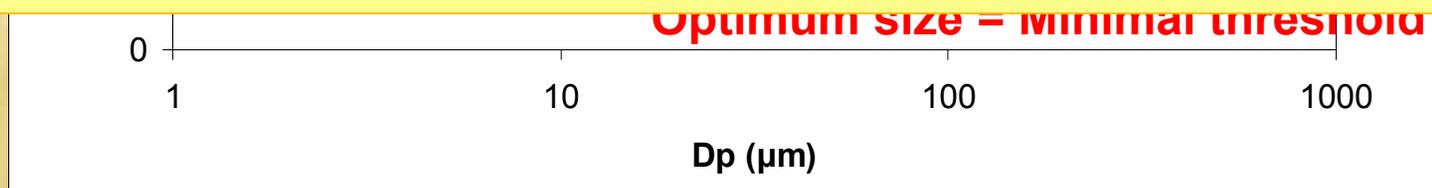


DUST EMISSIONS PROCESSES : THRESHOLD

A size-dependent erosion threshold

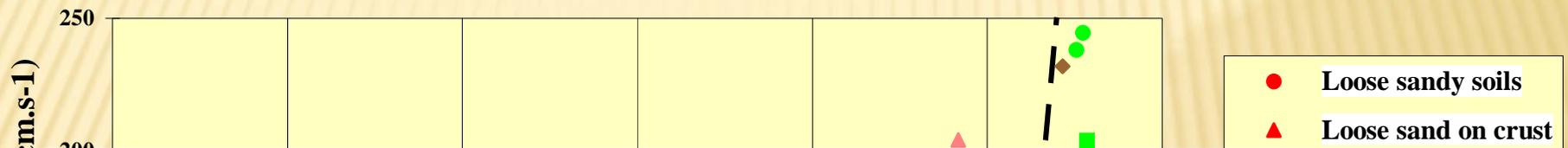


→ **The available parametrizations provide similar results and correctly simulate the erosion threshold as a function of soil grain size**



DUST EMISSIONS PROCESSES : THRESHOLD

Influence of surface roughness



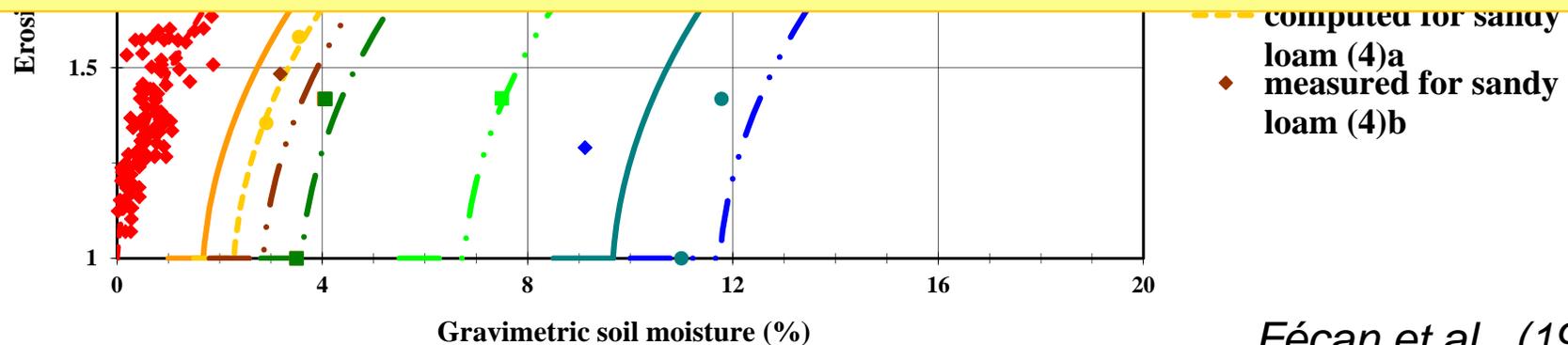
→ **Such parameterizations are OK for arid areas (solid obstacles; low roughness densities) but must be adapted for vegetated surface (arrangement; porosity, flexibility, etc ..)**

(Mac Kinnon et al., 2004)

DUST EMISSIONS PROCESSES : THRESHOLD

Influence of soil moisture

→ No additional experimental data sets than those used to establish the parametrization; seems OK in the field (Ishisuka et al. 2005)

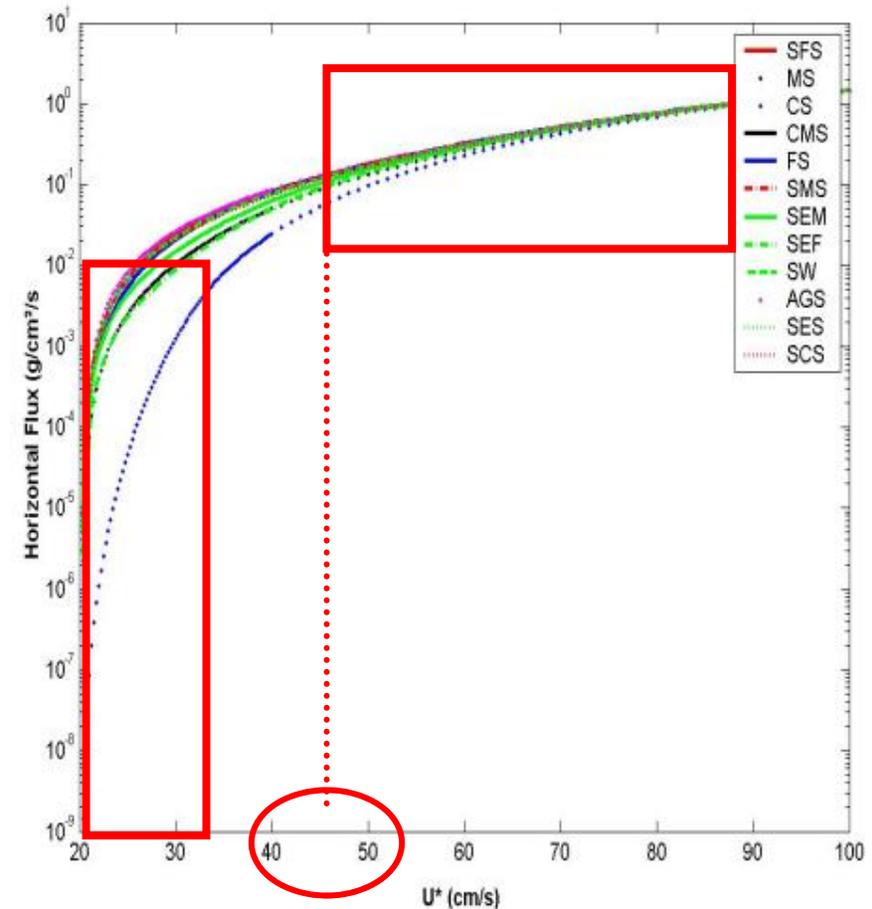
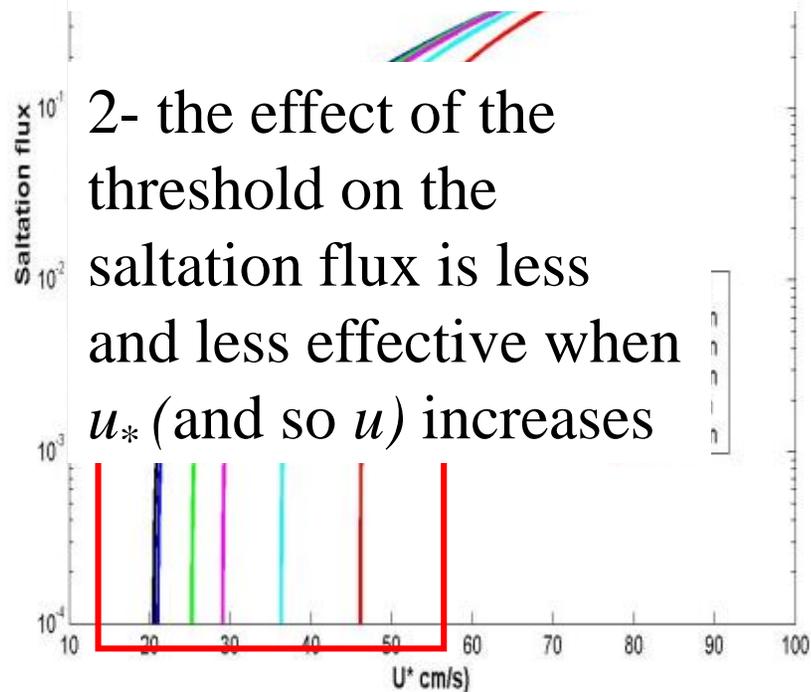


Fécan et al., (1999)

DUST EMISSIONS PROCESSES : SALTATION

1- When applied to natural soils, the influence of the soil grain size distribution on the threshold is minimized

2- the effect of the threshold on the saltation flux is less and less effective when u_* (and so u) increases

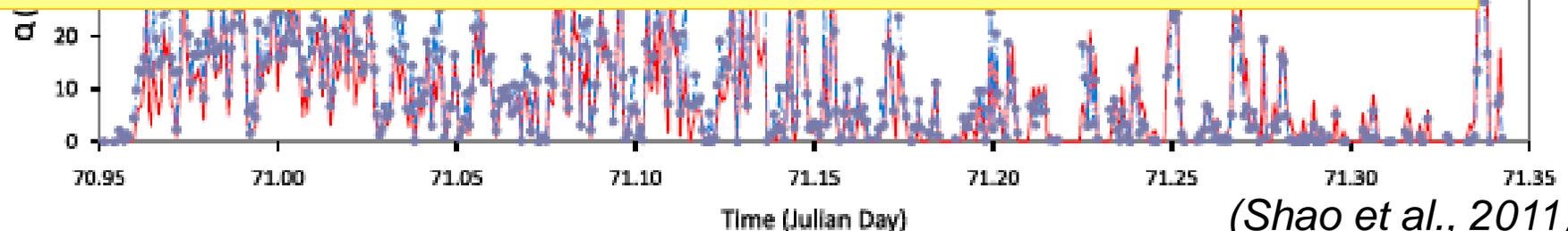


DUST EMISSIONS PROCESSES : SALTATION



$$Q = \int_{d_1}^{d_2} Q(d)p(d)\delta d$$

→ Physical understanding and available parametrization are OK; some tuning may be needed in the field (soil size distribution, fraction of uncovered surfaces, crust, etc ..)

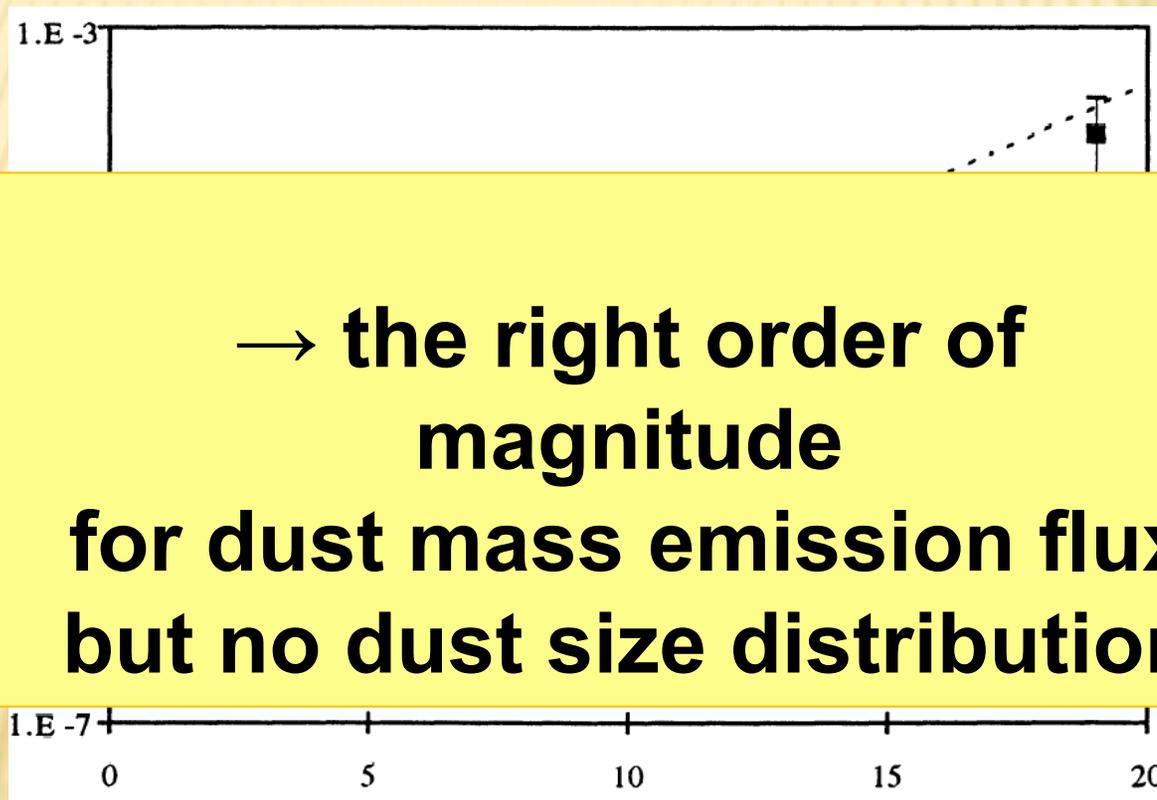


(Shao et al., 2011)

DUST EMISSIONS PROCESSES : SANDBLASTING

The soil clay content as a proxy of the amount of dust in the soil

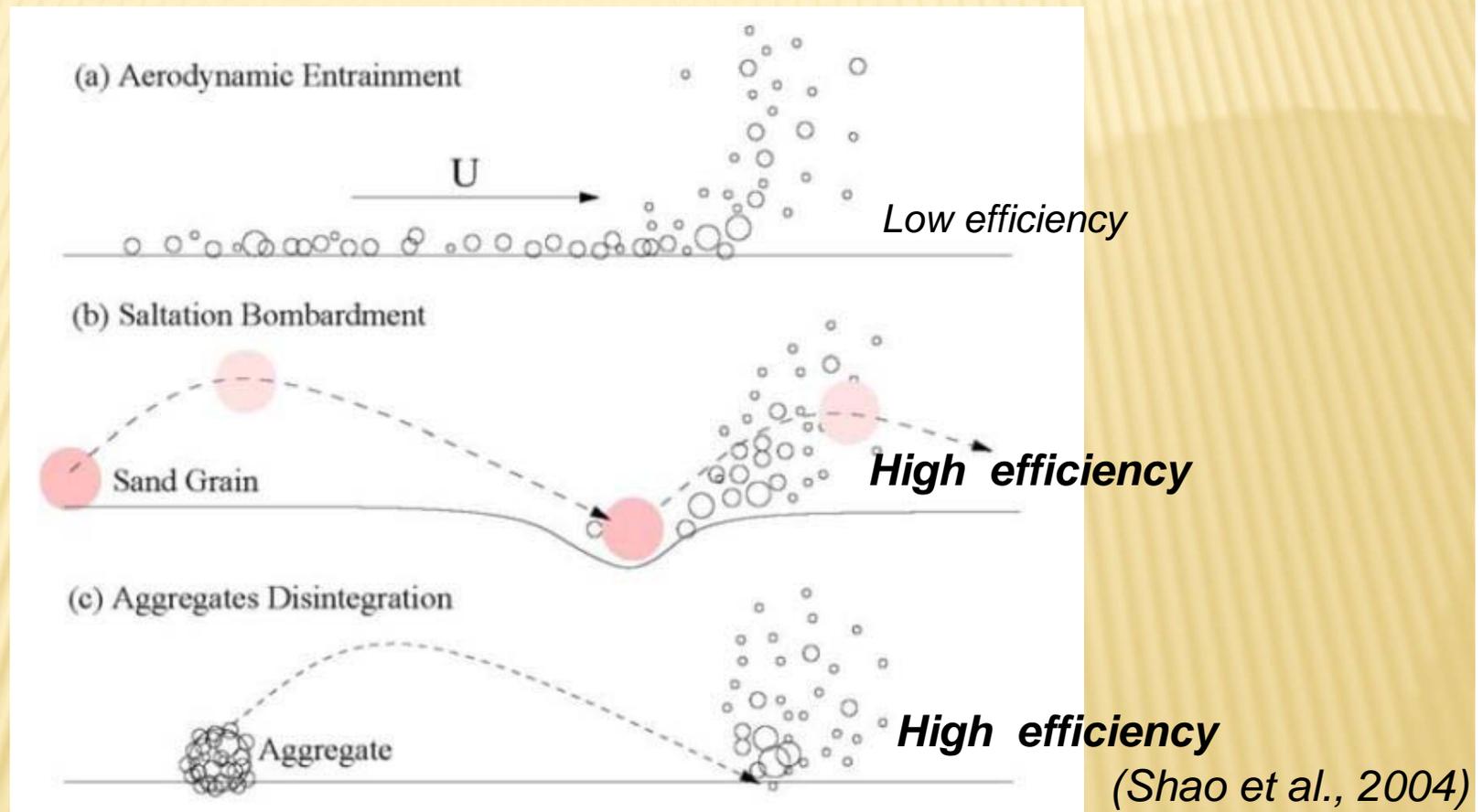
Sandblasting efficiency



→ the right order of magnitude for dust mass emission flux but no dust size distribution

(Marticorena and Bergametti, 1999)

DUST EMISSIONS PROCESSES : SANDBLASTING



= A balance between the kinetic energy of saltating grains and the cohesion of fine particles

DUST EMISSIONS PROCESSES : SANDBLASTING

⇒ Conceptual understanding

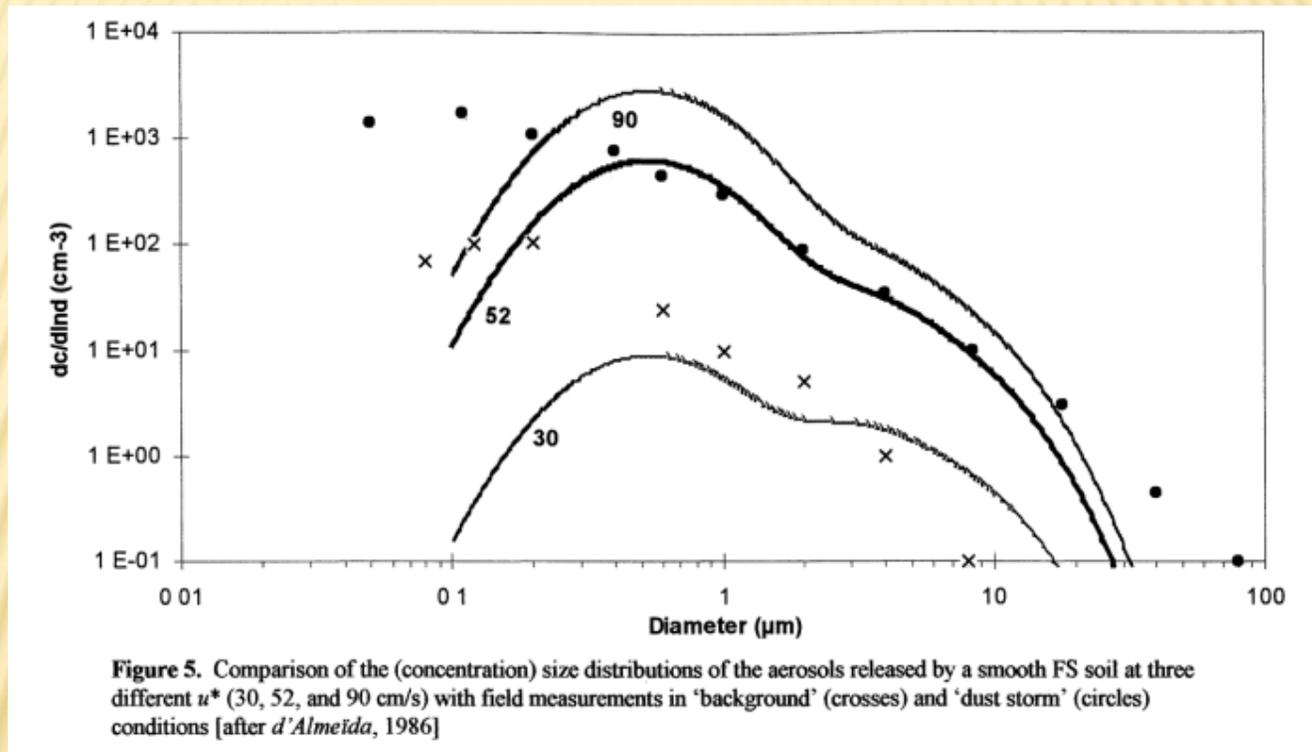
Kinetic energy provided

Binding energy
of the dust particles

- **all these models require input parameters that are difficult/impossible to measure**
- **Measurements of the size distribution of the emission fluxes could be used as a constrain**

DUST EMISSIONS PROCESSES : SANDBLASTING

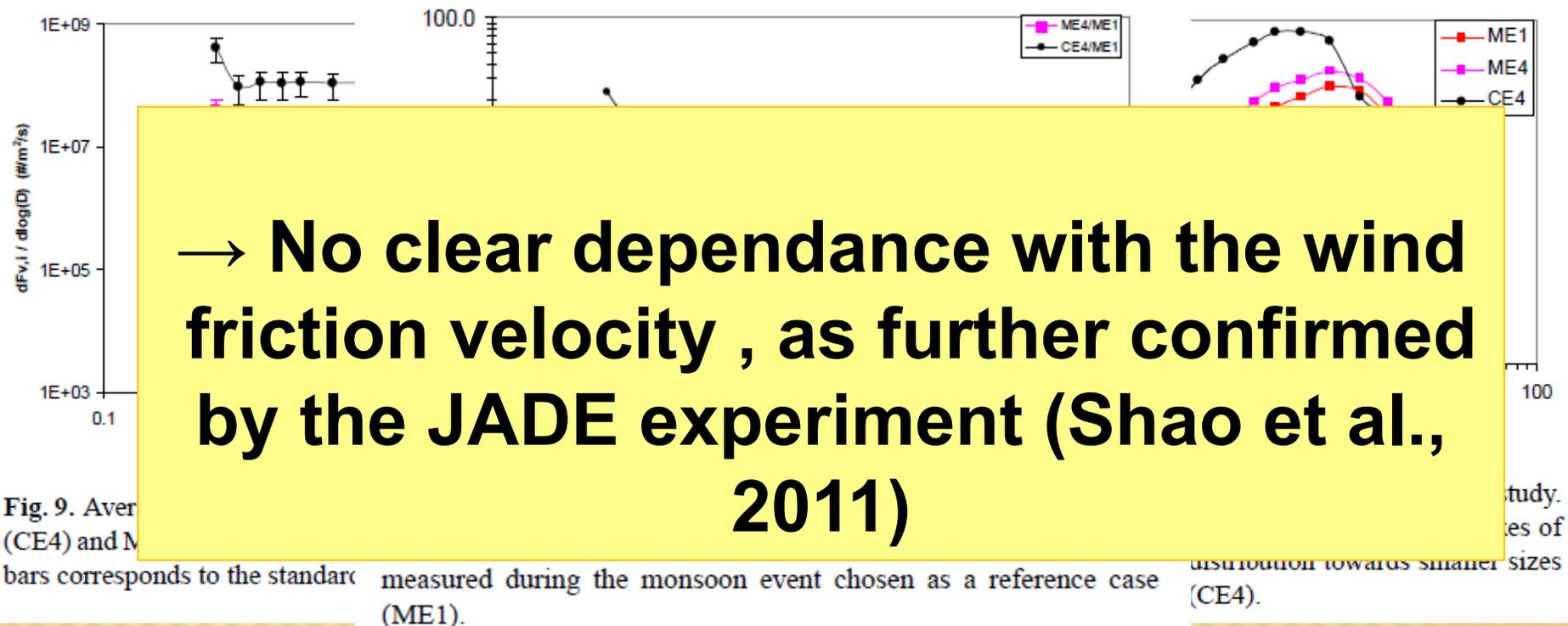
Simulated dust size distribution (Alfaro and Gomes, 2001)



-The proportion of the finest modes increases as wind friction velocity increases because of higher kinetic energy flux

DUST EMISSIONS PROCESSES : SANDBLASTING

Size-resolved dust emission fluxes (Sow et al., 2000)



DUST EMISSIONS PROCESSES : SANDBLASTING

Size-distribution of the dust emission fluxes (Sow et al., 2000)
 (log normal modes; gmd, = geometric mean diameter in μm)

	mode1			mode2			mode3		
	σ	gmd	%	σ	gmd	%	σ	gmd	%
ME1				1,7	4,9	42%	1,5	10,4	58
ME4				1,8	5,1	43%	1,5	10,4	57%
CE4	1,50	1,70	11%	1,7	5,1	77%	1,5	10,0	12%

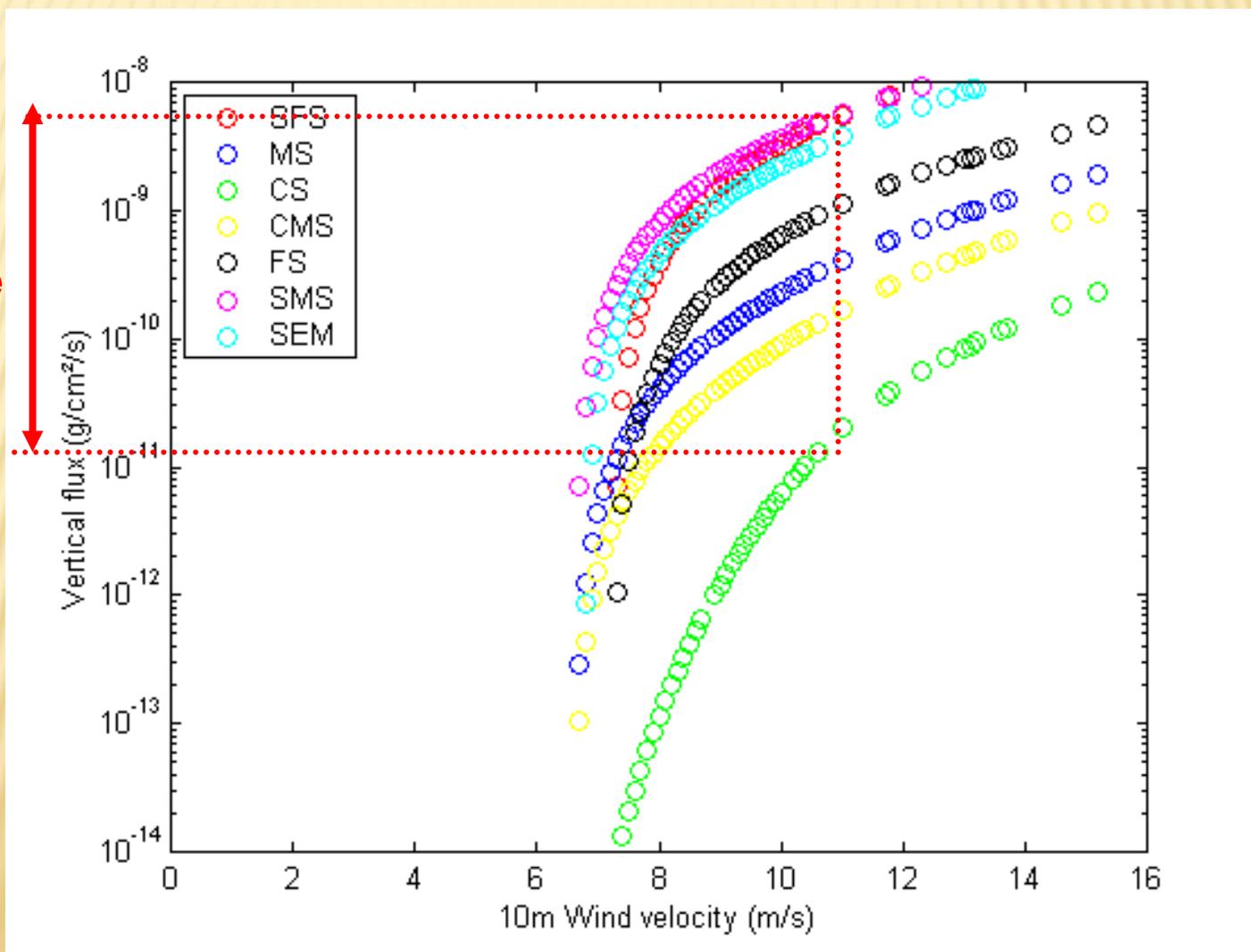
Wind tunnel 1.7 1.5 1.6 6.7 1.5 14.2 (Alfaro et al., 1998)

The two coarse modes are finer than in the wind tunnel exper

DUST EMISSIONS PROCESSES : SANDBLASTING

Sensitivity of the dust flux (Alfaro and Gomes, 2001) to the soil grain

3 orders
of
magnitude



DUST EMISSIONS PROCESSES : SANDBLASTING

Simulation of the size-distribution of emitted dust with a brittle fragmentation theory (Kok, 2011)

Input parameters

Propagation length

$\lambda = 12 \pm 1 \mu\text{m}$

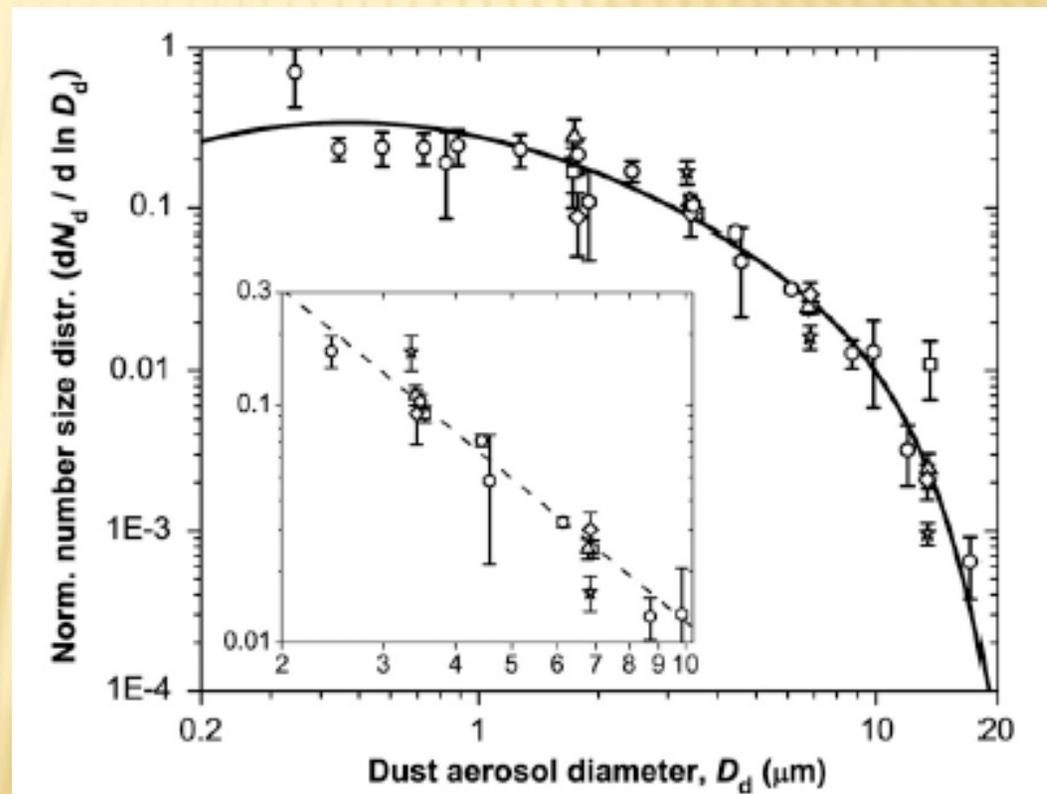
Soil texture

$D_s = 3.6 \mu\text{m}$

$\sigma_s = 3$

$C_N = 0.9539$

$C_V = 12.62$



Dust size distribution = power law

DUST EMISSIONS PROCESSES : SANDBLASTING

- **Soil properties seems to play a critical role in the sandblasting models ; to be further investigated (wind tunnel; in the field)**
- **Soil size distributions is not homogeneously measured (no standard)**
- **Measured dust size distribution must be enlarged (<0.3 μm ; >10 μm)**

DUST EMISSIONS PROCESSES

- **Dust emission fluxes can be reasonably well predicted if surface parameters (soils properties, surface roughness, vegetation/litter cover), surface winds and soil moisture are correctly measured/estimated**
- **Sand blasting models cannot predicted the size distribution of the emitted dust with a good confidence level.**

DUST EMISSIONS PARAMETERIZATIONS

- ✘ Formally, dust emission parameterizations used in 3-D models are in the form :

$$F \approx C \cdot U^3 (U - U_t)$$

- ✘ Simplifications : Prescribed preferential sources, erodability index ($C = \sum c_i$) , ...
- ✘ Input parameters
 - + Surface properties: aeolian roughness; soil size-distribution
 - + Meteorological parameters : Wind velocity ; soil moisture or precipitation

INPUT PARAMETER : SURFACE ROUGHNESS

✘ Mapping from satellite surface products

→ **Surface satellite products (BRDF, radar backscatter coefficient) provides very good estimation of the aeolian roughness length and thus of the erosion threshold**

→ **Today different maps are available at the global scale (*i.e.*, Prigent et al., submitted to AMT-D; MODIS BRDF, E. Vermote))**

INPUT PARAMETER : SURFACE ROUGHNESS

Dust sources are clearly emerging

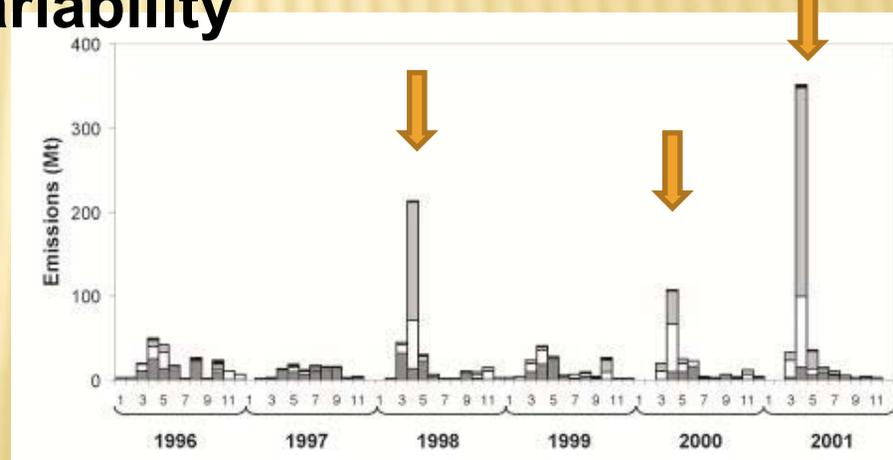
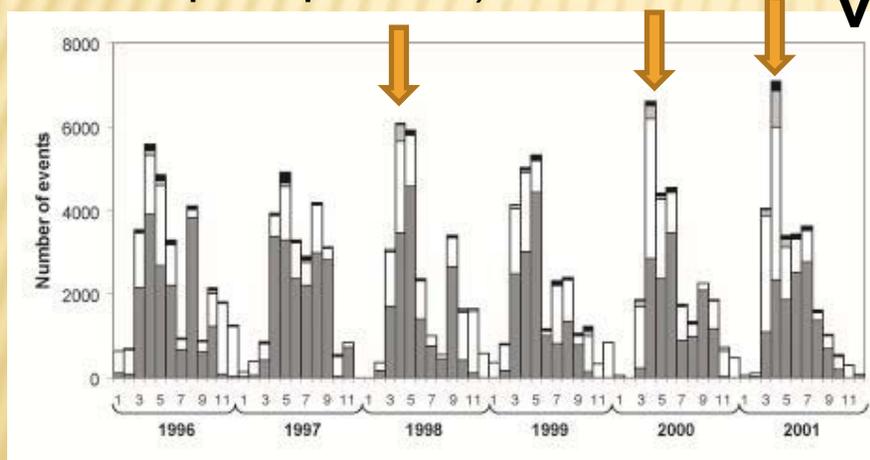
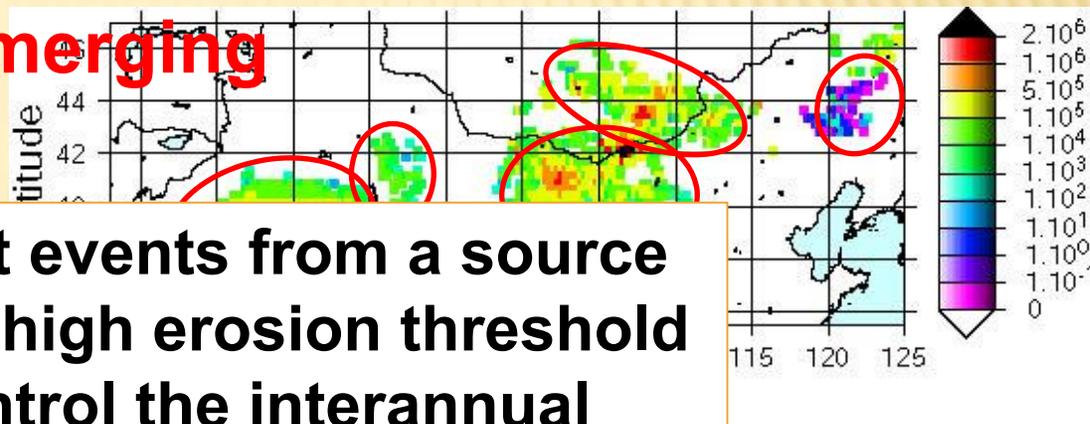
Simulated dust emissions

242±131 Mt

- 1996-2001

(ERA40 surface wind precipitation)

3 dust events from a source with a high erosion threshold control the interannual variability



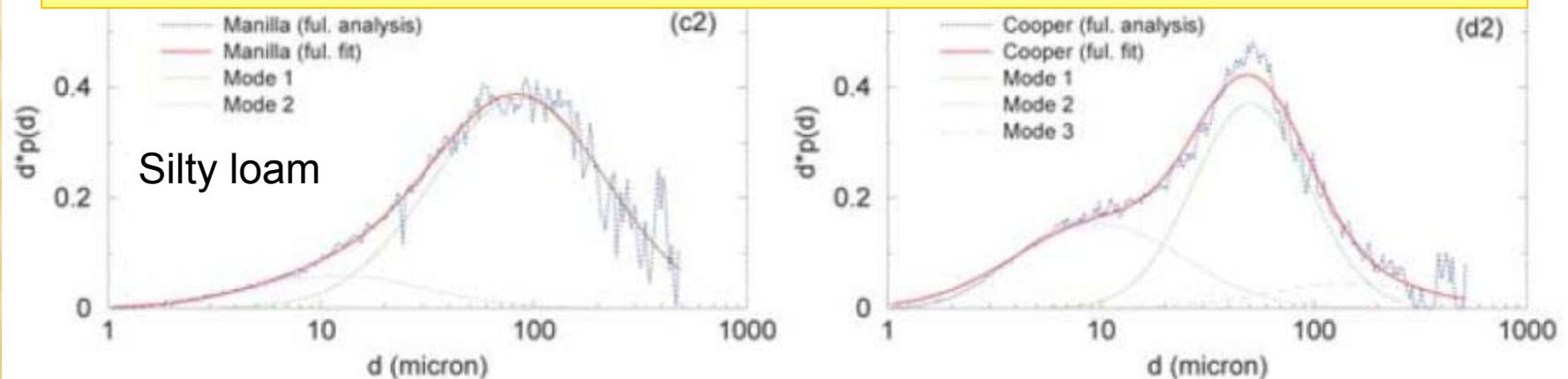
■ Taklimakan desert □ Northern deserts □ Gobi desert ■ Other arid areas

(Laurent et al., 2006)

INPUT PARAMETER : SOIL SIZE DISTRIBUTION

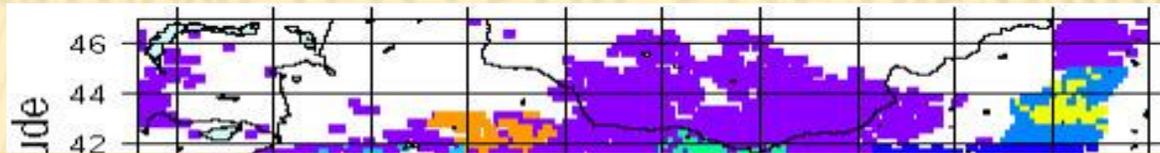
✘ Input for saltation flux computation =

→ **Undisturbed size distribution significantly differs from the disturbed one**



(Courtesy of Y. Shao)

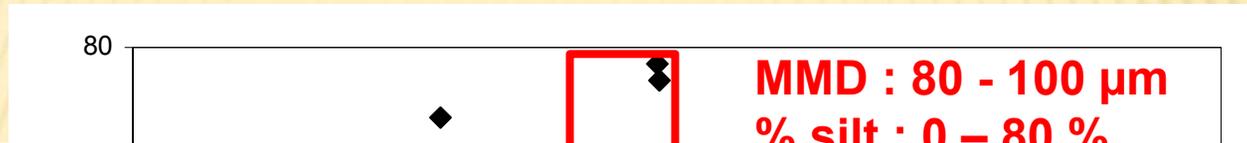
INPUT PARAMETER : SOIL SIZE DISTRIBUTION



→ In East Asia, undisturbed soil size are significantly different from one desert to the other and relatively homogeneous in a given desert

Ulan Buh and Badain Jaran (N=10)		97	1.30	52	316	1.59	48	3.4	8.6	88.2
Tengger and Kubqi (N=9)		120	1.48	72	322	1.29	28	2.6	7.3	90.7
Mu Us (N=8)		99	1.17	35	330	1.37	65	1.6	7.7	90.2
Horqin (N=23)		315	1.29	100	-	-	-	1.6	7.7	90.2
East of Xinjiang area (N=4)		90	1.24	29	293	1.66	71	9.9	34.7	55.3
Hexi Corridor (N=10)		97	1.26	40	386	1.59	60	4.8	14.8	80.6
Gurban Tunggut (n=3)		94	1.12	36	170	1.69	64	3.6	13.5	82.0

INPUT PARAMETER : SOIL SIZE DISTRIBUTION



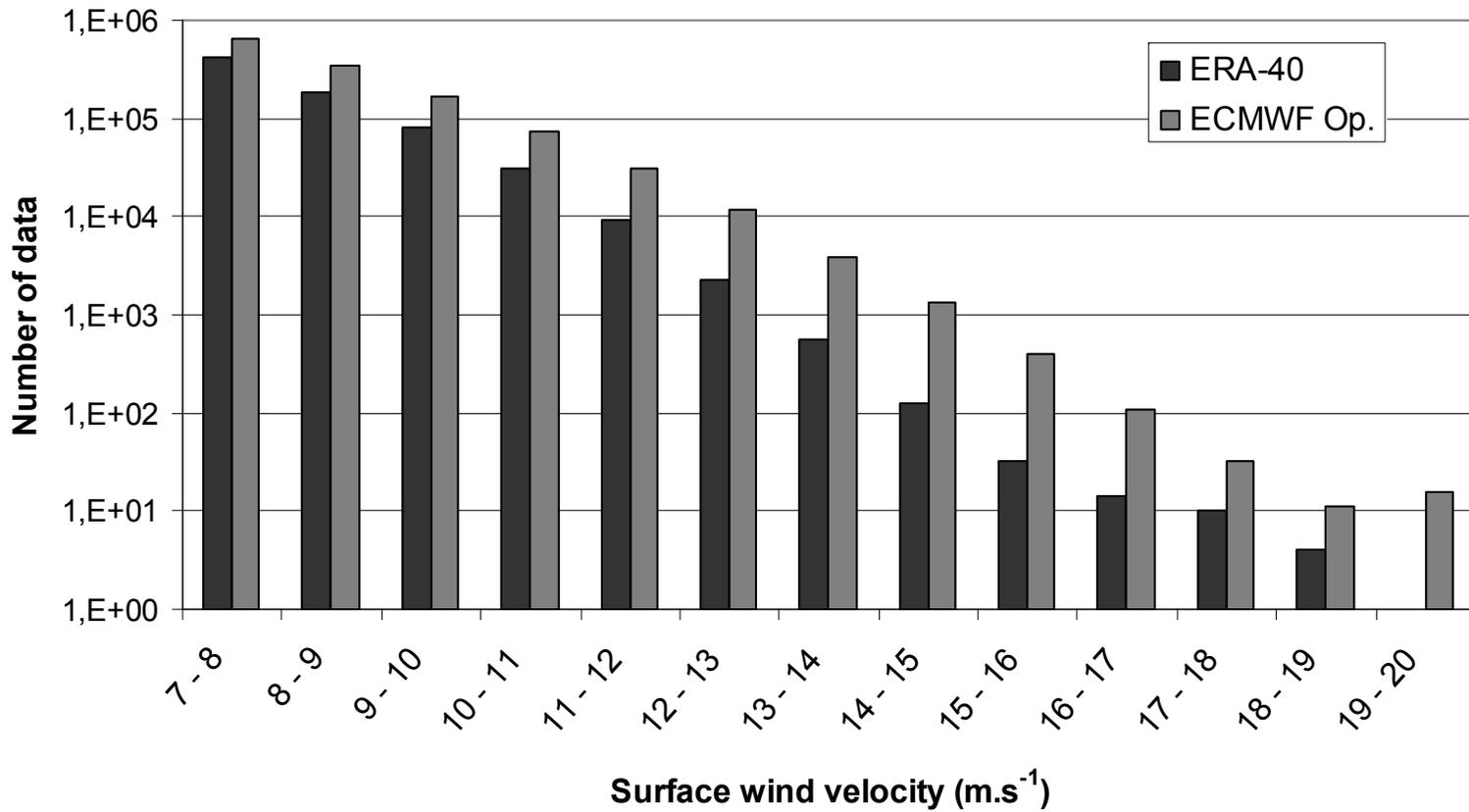
→ **No unambiguous relationship between soil texture and undisturbed soil size distribution**

→ **Additional samples/analyses are needed for other deserts to establish links with soil texture (available at global scale)**

(Laurent et al., 2006)

INPUT PARAMETER : WIND VELOCITY

Annual mean emissions



Lau
Sch

3
4
7 M

DUST SOURCES SIMULATION : VALIDATION ?

- **Difficult to validate simulated dust emissions due to uncertainties on aerosol satellite products over bright surfaces**
- **In some regions horizontal visibility from meteorological stations can be used**
(when a sufficient number of stations and data is available)
- **Quantification of the relevance of the dust emissions through results of 3-D regional simulations ?**

REGIONAL SIMULATIONS WITH CHIMERE-DUST

Dust Emission :

⇒ Dust emissions fluxes (Marticorena and Bergametti, 1995); Surface data base (Marticorena et al., 1997)

⇒ Associated size-distribution (Alfaro and Gomes, 2001)

⇒ distributed on 20 log. bins

Domain : 10S-60N, 90W-90E

North Atlantic, North Africa, Arabian Peninsula

Model outputs

-Dust 4D fields in $\mu\text{g}/\text{m}^3$ for each bin
-Optical thickness, deposition fluxes

Simulation domains:

- Horizontally: (1x1 degrees)
- Vertical mesh 15 to levels (up to 200hPa)

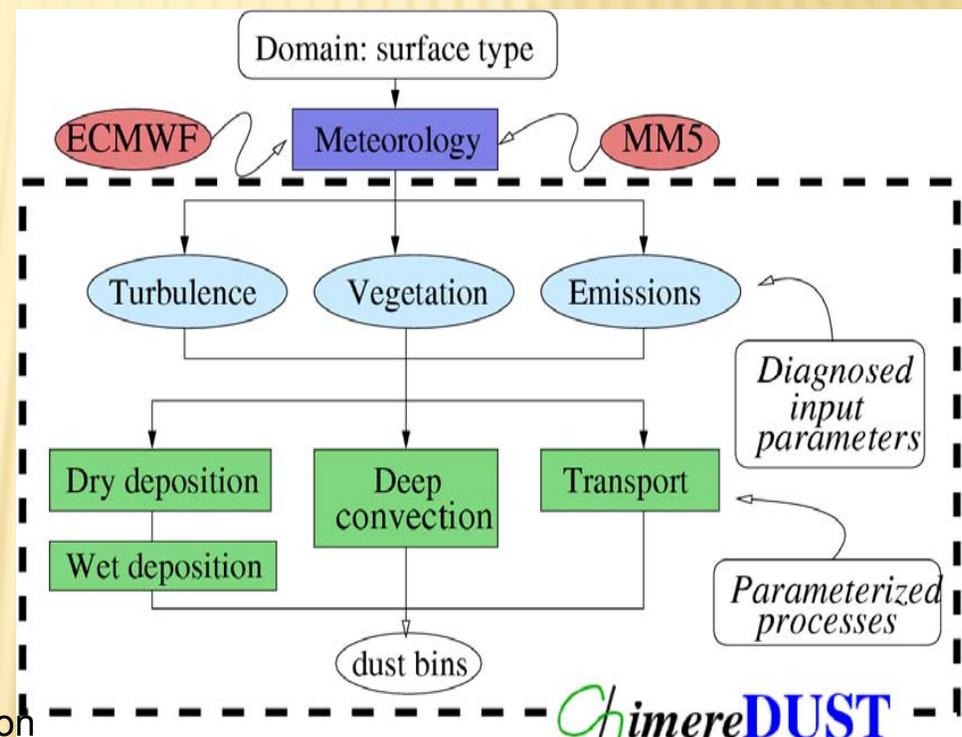
Meteorological forcing

-ECMWF forecast (First guess) + empirical correction of surface winds in the Bodélé Depression

Aerosol Optical Depth @550nm

Refractive Index = $1.5 - 0.005i$; (Moulin et al., 2001)

A Chemistry and Transport Model
with no chemistry but dust ...

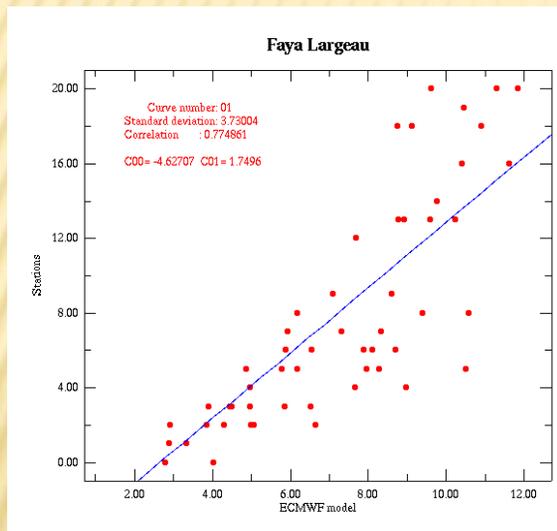


INPUT PARAMETER : WIND VELOCITY

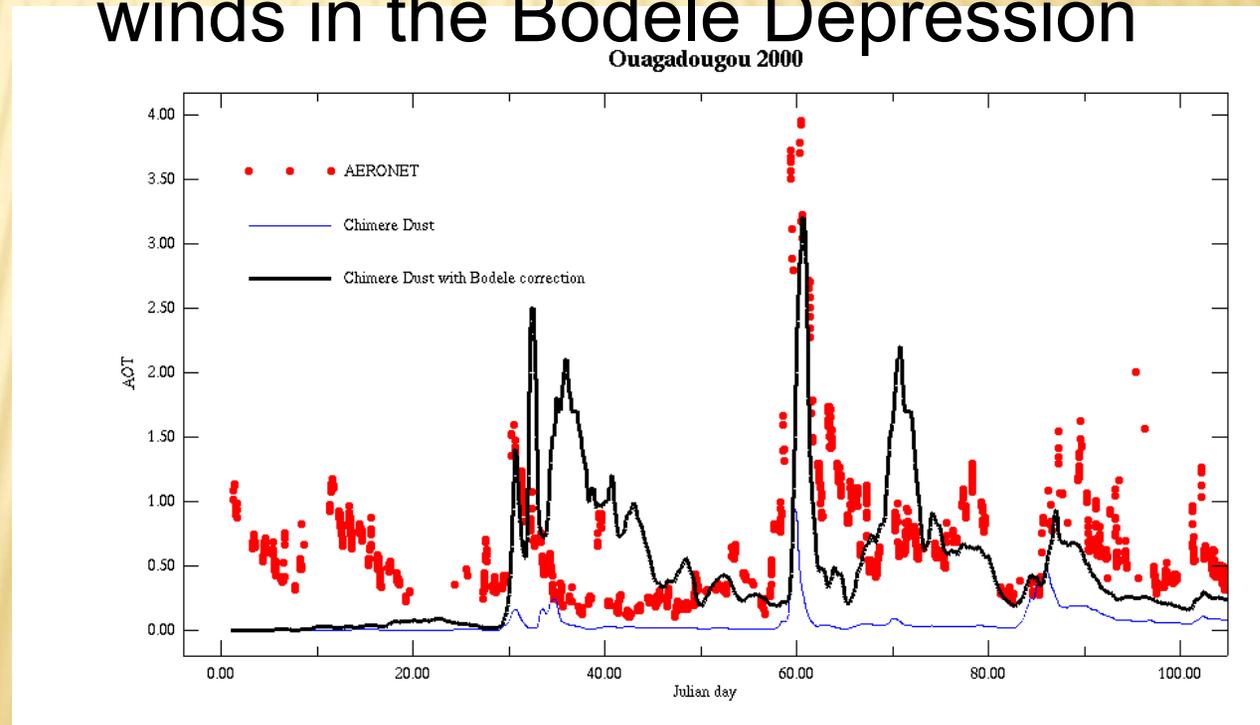
Simulation of the dust content
over the Sahel with CHIMERE-Dust

:

Impact of the correction of surface
winds in the Bodele Depression



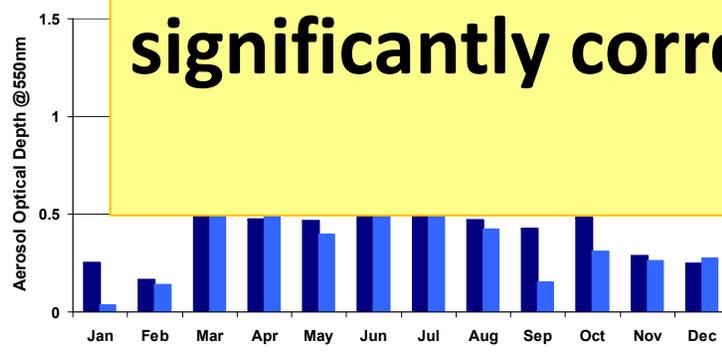
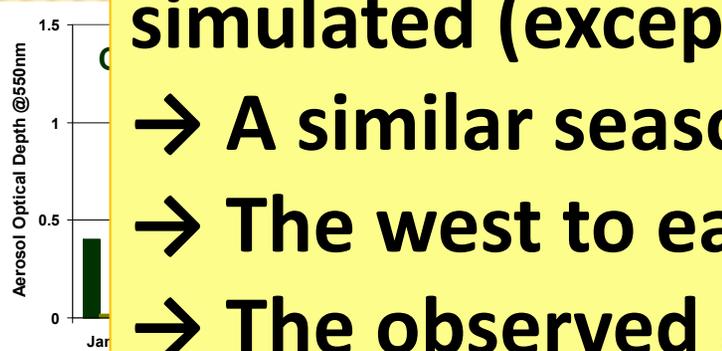
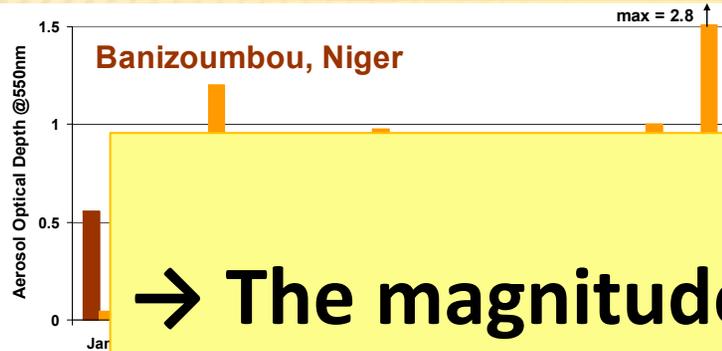
**Correction based on
measurements in
Faya-Largeau
15N-19N and 15E-
20E**



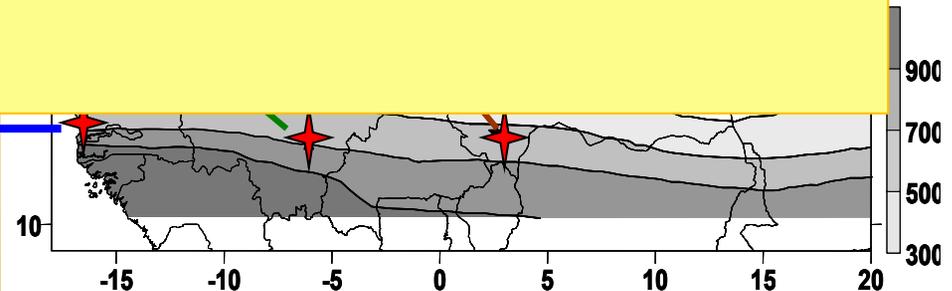
REGIONAL SIMULATIONS WITH CHIMERE-DUST



ChimereDUST



- The magnitude of the observed AOD is well simulated (except jan, dec)
- A similar seasonal cycle at the three stations
- The west to east gradient is retrieved
- The observed and simulated AODs are significantly correlated ($n=36$; $r=0.53$)

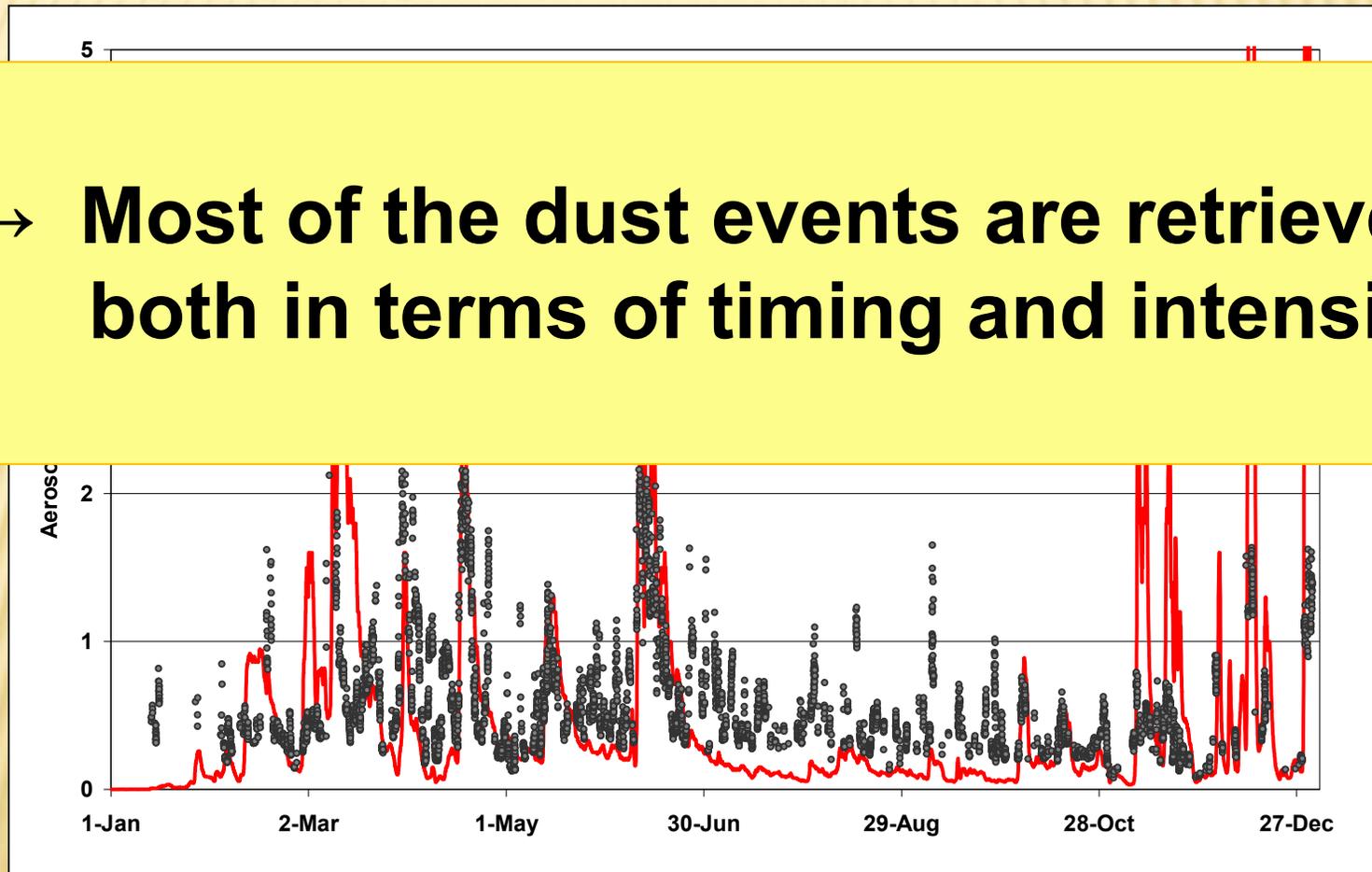


(Observed : dark color; Simulated : light color)

REGIONAL SIMULATIONS WITH CHIMERE-DUST

Hourly measured and simulated aerosol optical depth

→ **Most of the dust events are retrieved both in terms of timing and intensity**



(Level 2 AODs with $\alpha > 0.4$)

(Schmechtig et al., 2011)

REGIONAL SIMULATIONS WITH CHIMERE-DUST

Daily measured and simulated surface concentration

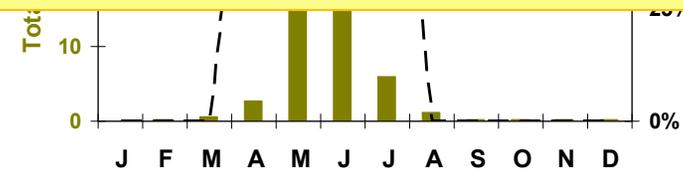
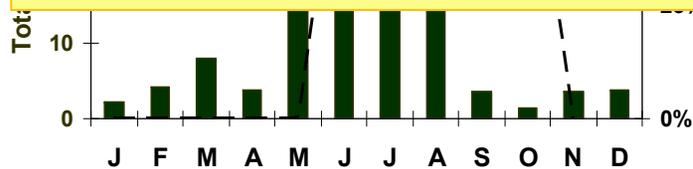
- The order of magnitude of the surface concentrations is retrieved**
- The seasonal cycle is well reproduced**
- The level of agreement with observations is similar than PM concentration in air quality models (NME = 75%; NMB = -36 %)**

REGIONAL SIMULATIONS WITH CHIMERE-DUST

Year 2006	Total deposition ($\mu\text{g}\cdot\text{m}^{-2}$)	
	Measured	Simulated
MD	83.8	59.6

⇒ Annual total deposition fluxes are reasonable but underestimated

- Underestimation of the dry deposition ?
- Bias in the size distribution ?
- Significant bias due to precipitation spatial and temporal distribution



CONCLUSION (1/3)

- ✘ **Dust emission fluxes can be reasonably well predicted based on available parameterizations provided surface parameters are correctly estimated**
 - + **Surface roughness can be mapped based on satellite products**
 - + **Soil properties must be derived/calibrated from measurements**

- ✘ **Emitted dust size distribution of the emitted dust cannot be modelled today with a good confidence level.**
 - + **Additional process studies and field measurements are needed**
 - + **It can be assigned from available field measurements (AMMA, IADP)**

SIMPLIFIED DUST EMISSION SCHEME

Structure of the dust emission model Shao et al., 2004

1: Input
 $u_*; p_f(d), p_m(d), P, \theta, \theta_r, \lambda, \sigma$

2: Ideal threshold friction velocity
 $u_{*r0}(d) = \sqrt{a_1 \left(\frac{\rho_p}{\rho_a} g d + \frac{a_2}{\rho_a d} \right)}$
 $a_1 = 0.0123$
 $a_2 = 3 \times 10^{-4} \text{ kg s}^{-2}$

3: Real threshold friction velocity
 $u_{*r}(d; \lambda, \theta, s_c, s_r, \dots) = u_{*r0}(d) f_\lambda f_\theta f_{sc} f_{sr} \dots$
 $f_\lambda = (1 - m_r \sigma_r \lambda)^{0.5} (1 + m_r \beta_r \lambda)^{0.5}$
 $\beta_r \approx 202; m_r \approx 0.16, \sigma_r \approx 0.45$
 $f_\theta = [1 + A(\theta - \theta_r)^b]^{1/2}$
 $f_{sc} = f(s_c); f_{sr} = f(s_r)$

4: Saltation flux for particle size d
 $Q(d) = \begin{cases} c_o \sigma \frac{\rho_a}{g} u_*^3 \left(1 - \frac{u_{*r}^2}{u_*^2} \right) & u_* \geq u_{*r} \\ 0 & u_* < u_{*r} \end{cases}$

5: Saltation flux Q
 $Q = \int_{d_1}^{d_2} Q(d) p(d) \delta d$
 $p(d) = (1 - \gamma) p_m(d) + \gamma p_f(d)$

6: Dust emission for i^{th} bin
 $F(d_i, d_s) = c_y \eta_{\beta} [(1 - \gamma) + \gamma \sigma_p] (1 + \sigma_m) g \frac{Q}{u_*^2}$
 $\sigma_m = 12 u_*^2 \frac{\rho_b}{P} \left(1 + 14 u_* \sqrt{\frac{\rho_p}{P}} \right)$
 $\sigma_p = \frac{p_m(d)}{p_f(d)}$
 $\gamma = \exp \left[-\frac{(u_* - u_{*r})^3}{\dots} \right]; \eta_{\beta} = \int_i p_f(d) \delta d$
 $F = \int_i F(d_i, d) p(d) \delta d$

7: Dust emission
 $F = \sum_i F_i$

Simplified Soil types

A = F/Q = f(clay content)

Prescribed updated size-distribution at emission

(Marticorena and Bergametti, 1995)

(Sow et al., 2009; Shao et al., 2011)

CONCLUSION (2/3)

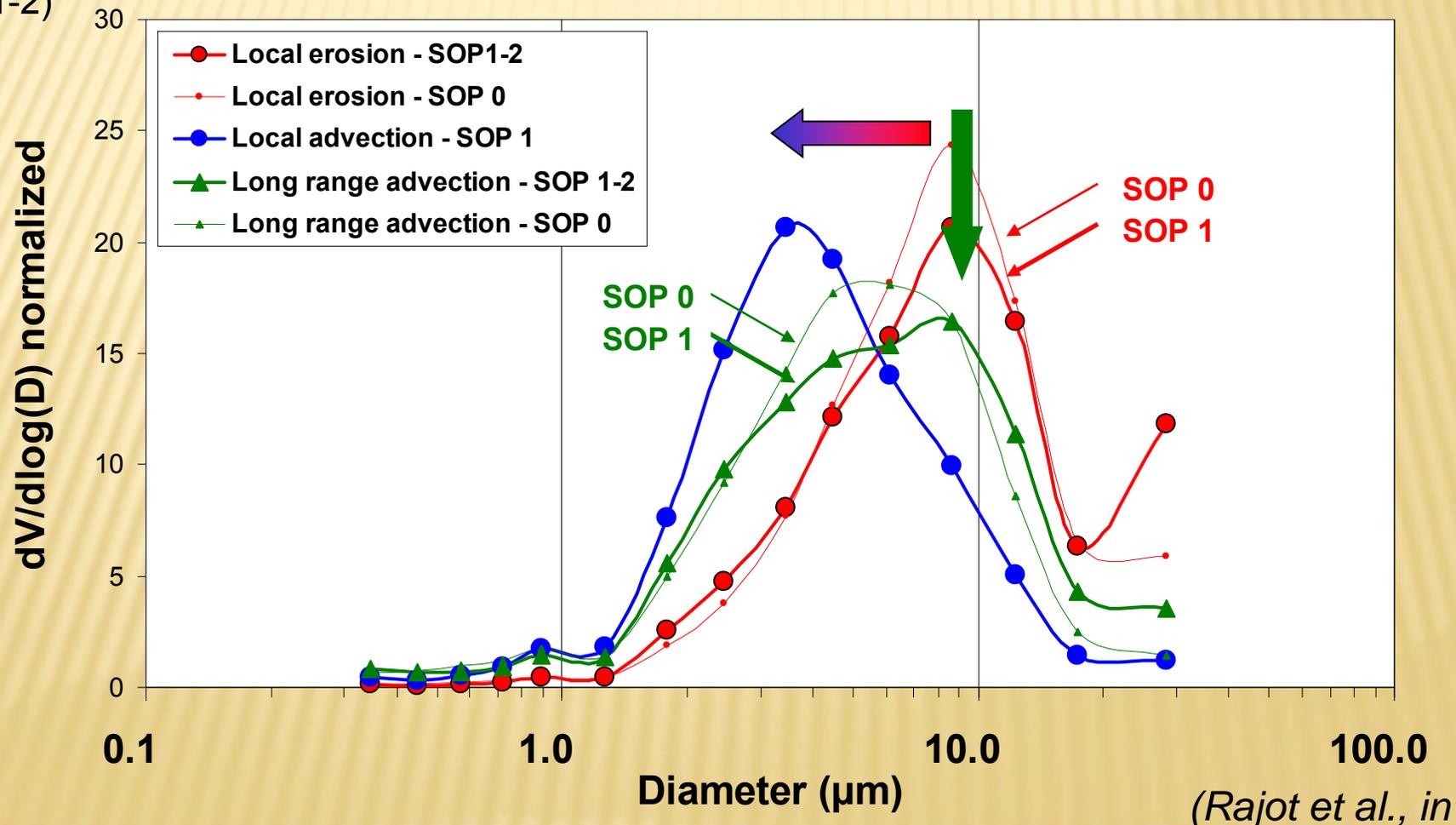
- × **Surface wind velocity is the most critical parameter for dust emission simulations**
- × The capability of meteorological models to provide **realistic surface winds** must be **questioned and evaluated**
- × **Bias** in the simulated surface wind must be **corrected** ; it **cannot be compensated** by a tuning of the surface properties
- × **Soil moisture** is also important in defining source regions and suffers from even larger uncertainties

CONCLUSION (3/3)

- ✘ Satellite aerosol products should be used to constrain the dust events frequency separately from the dust content (a constrain on the erosion threshold/modelled wind velocity)
- ✘ Other uses of satellite products should be investigated (evaluation displacement velocities of dust plumes; gradient of deposition using angstrom coefficients, ...)
- ✘ **The only way to reduce the uncertainties on the dust emissions is to close the mass budget**
 - + **Deposition** networks are needed
 - + **Data on dust size distribution** from the recent intensive field campaign (large size range) must be compiled to extract trends and patterns to test the 3-Dmodels

DUST SIZE DISTRIBUTION MEASURED IN THE SAHEL

- ✓ Shift between local erosion and local advection (decrease of coarse modes)
- ✓ Coarse mode within the long range transport
- ✓ Local erosion and Long range advection similar within dry (SOP 0) and wet season (SOP 1-2)





Thank you for your attention !